

Simulation of the transport phenomena in the Horstberg geothermal system

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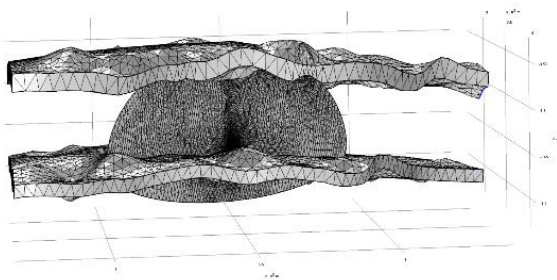
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Motivation



- The Horstberg geothermal reservoir in the North German Basin (80 km north of Hannover)

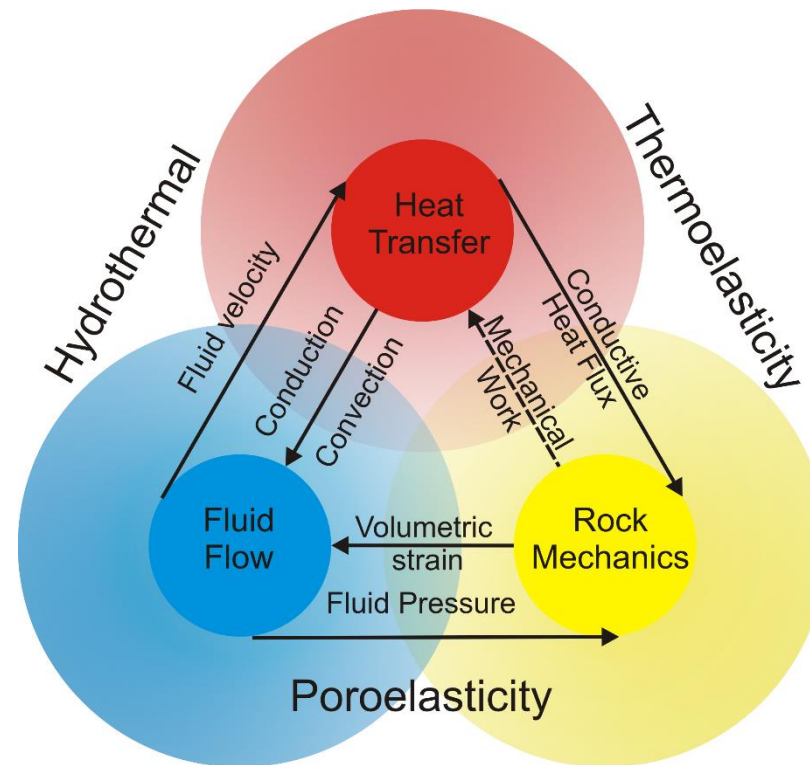
- Reservoir depth: 3650 m TVD
- Reservoir temperature: 144 °C
- Reservoir pressure: 584 bar
- Shmin: 650 to 680 bar



- An induced hydraulic fracture hydraulically connects two sandstone layers
- Area of induced fracture: 196,000 [m²]

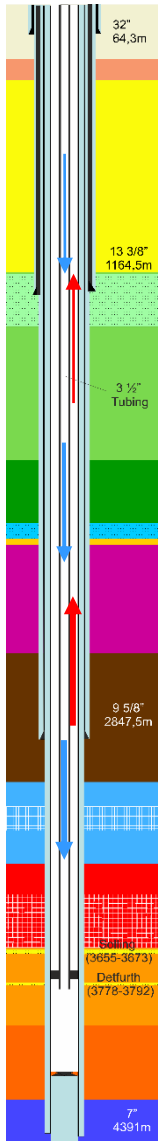
Research Question

- How to simulate transport phenomena in a geothermal system

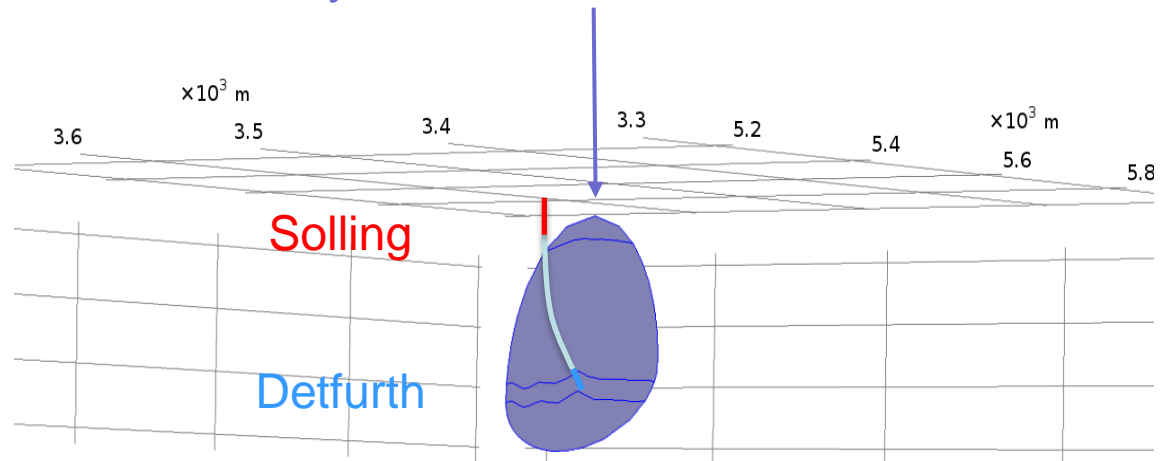


(Hassanzadegan, 2012 PhD thesis)

Horstberg Wellbore



- A monoborehole concept was applied in a multilayered sandstone reservoir
- Production through annulus from **Solling Formation** and injection through tubing into **Detfurth Formation**
- Induced hydraulic fracture



- Laminar Flow : velocity and pressure fields

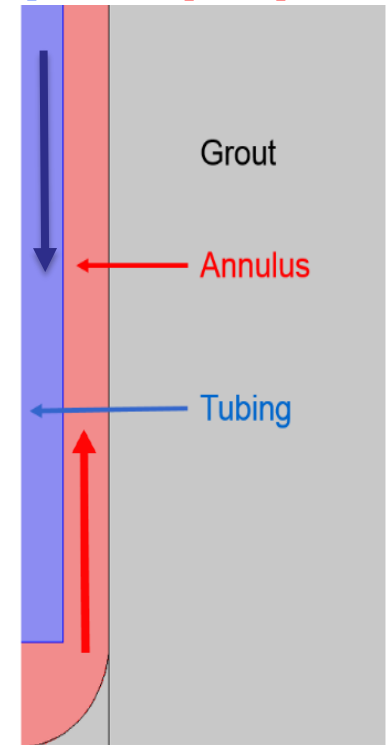
- Conservation of mass

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0$$

- Conservation of momentum

$$\rho \frac{\partial \mathbf{u}}{\partial t} + \rho (\mathbf{u} \cdot \nabla) \mathbf{u} = \nabla \cdot [-p\mathbf{I} + \boldsymbol{\tau}]$$

Q=0.1 [lit/s] P=1 [atm]

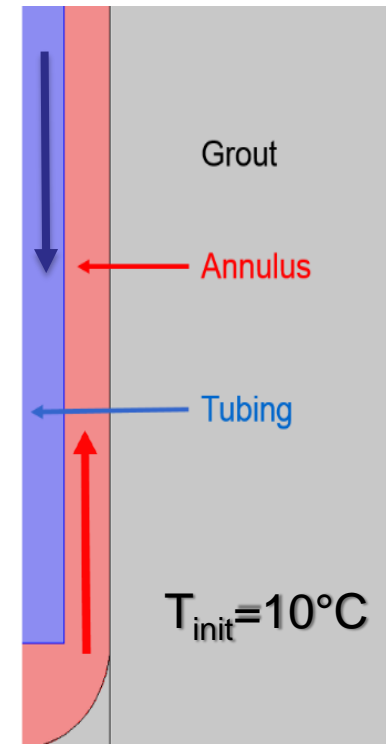


- Heat Transfer in Fluid: temperature field

$$\rho c_p \frac{\partial T}{\partial t} + \rho c_p \mathbf{u} \cdot \nabla T = \nabla \cdot (\lambda_f \nabla T) + \boldsymbol{\tau} : \nabla \mathbf{u}$$

T	Temperature
c_p	Heat capacity
λ	Thermal conductivity
\mathbf{u}	fluid velocity vector
ρ	density
$\boldsymbol{\tau}$	viscous stress tensor

$T=50^\circ\text{C}$ $T=?$

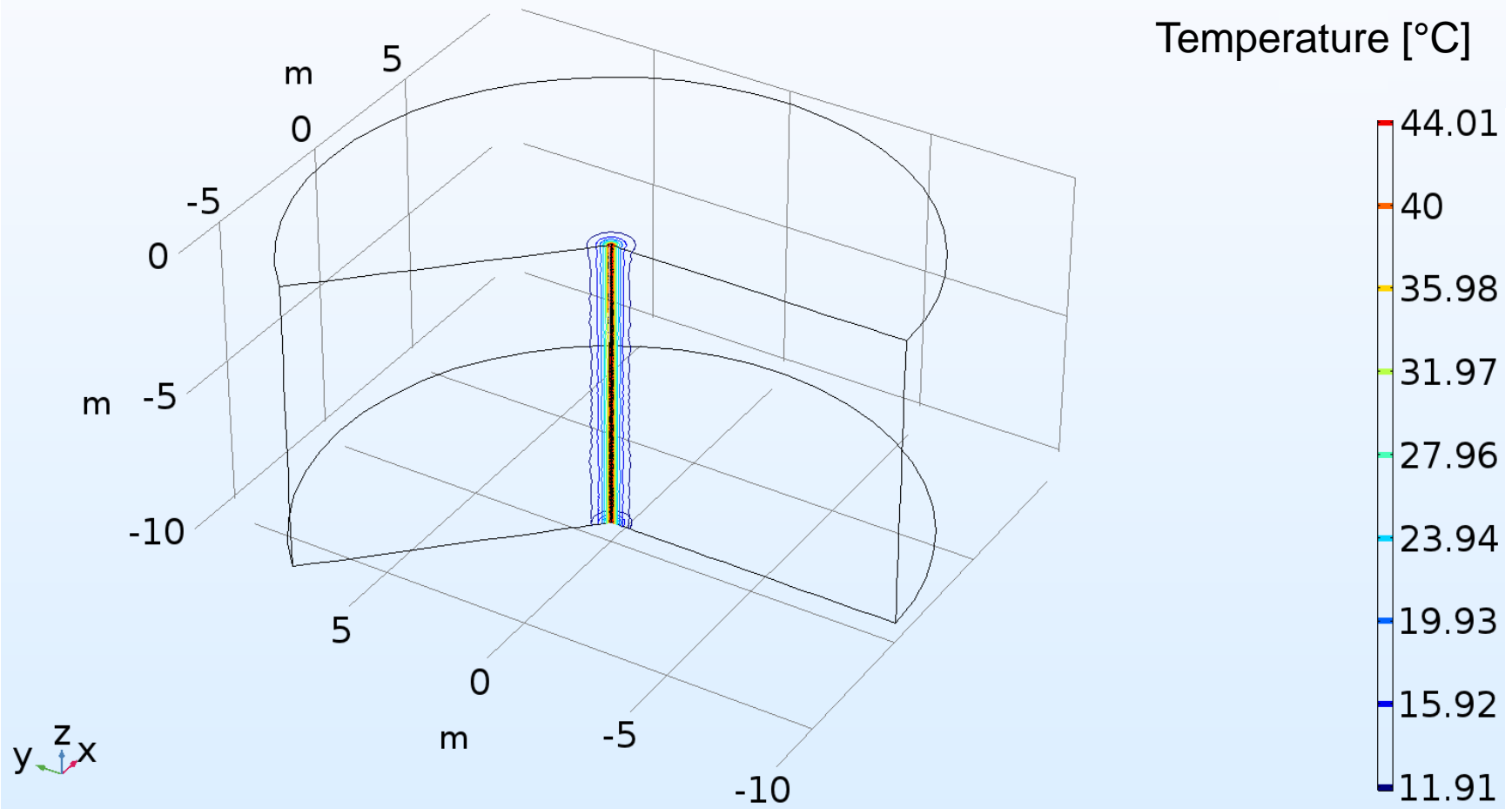


- The fluid flow was first solved in a stationary study and the resulting pressure and velocity fields were used in heat transfer study

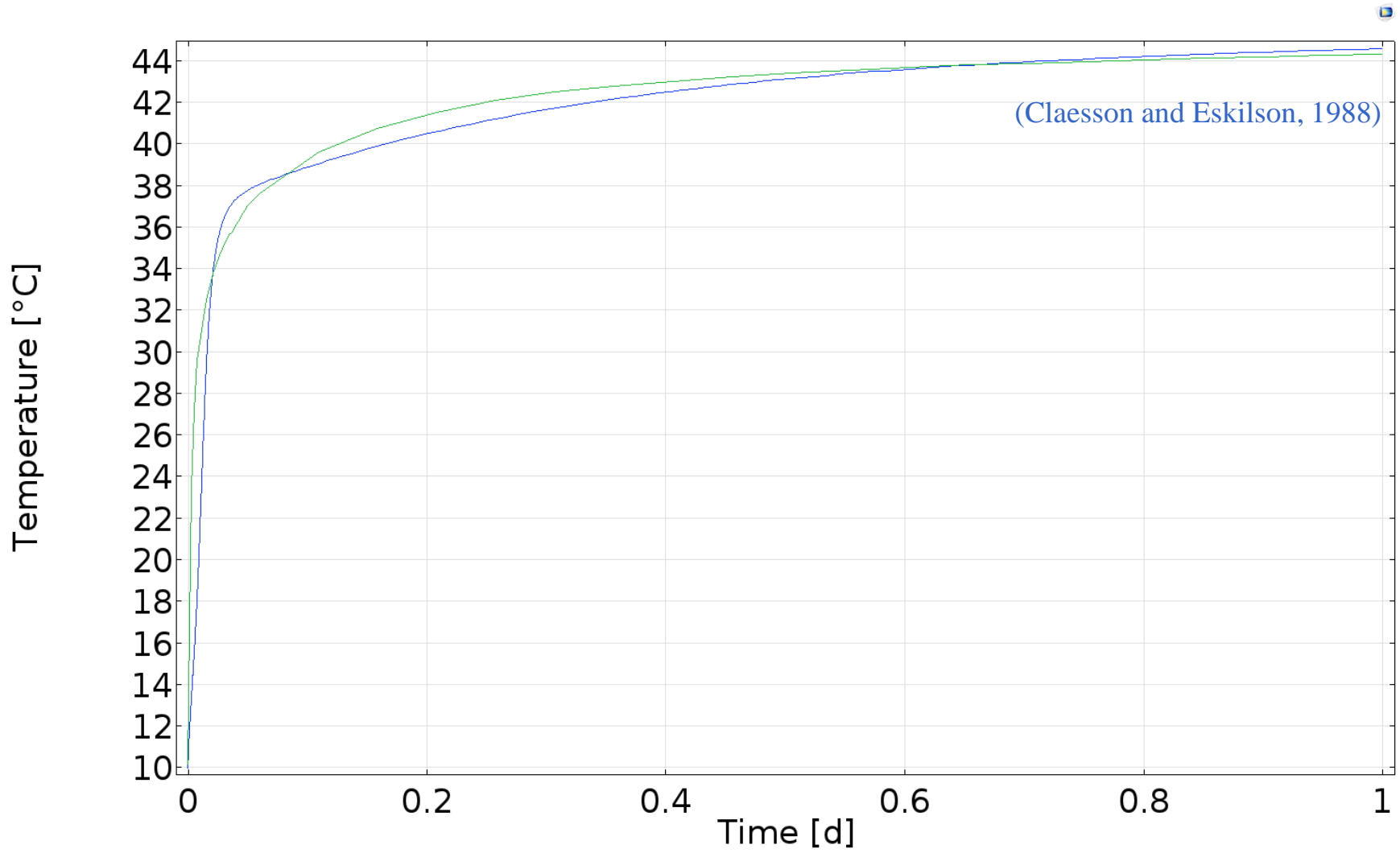
Verification Model

Axial Symmetry

Time=1 d Contour: Temperature (degC)



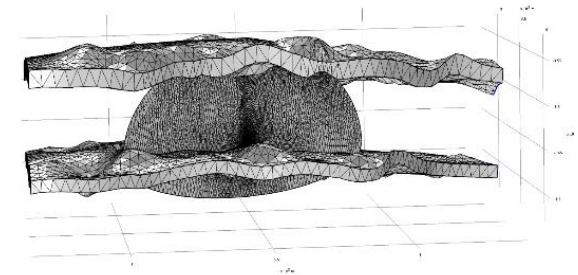
Verification Model



- Transient poroelastic equation:

$$\frac{\rho}{M} \frac{\partial p}{\partial t} + \nabla \cdot \rho \left[-\frac{k}{\mu} (\nabla p + \rho g z) \right] = -\rho \alpha \frac{\partial \varepsilon_v}{\partial t}$$

(Biot, 1962)



Storage Mass conservation Sink/Source term

- Constitutive poroelastic equations:

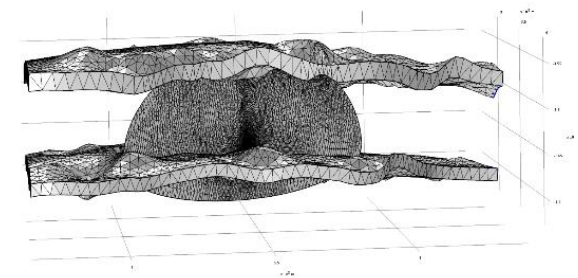
$$\boldsymbol{\sigma} = \mathbf{C} \boldsymbol{\varepsilon} - \alpha p \mathbf{I}$$

$$p = M(\zeta - \alpha \varepsilon_v)$$

σ	Stress
M	Biot Modulus
ζ	Fluid mass content increment
\mathbf{C}	Stiffness matrix
α	Biot coefficient ($0 \leq \alpha \leq 1$)
ε	Bulk strain
ε_v	Volumetric strain
\mathbf{I}	Unit tensor
p	Pore pressure

- heat transfer in porous media

$$(\rho C_p)_{eff} \frac{\partial T}{\partial t} + \nabla \cdot (\rho C_p \mathbf{u} T) + \nabla \cdot \mathbf{q} = Q_h$$

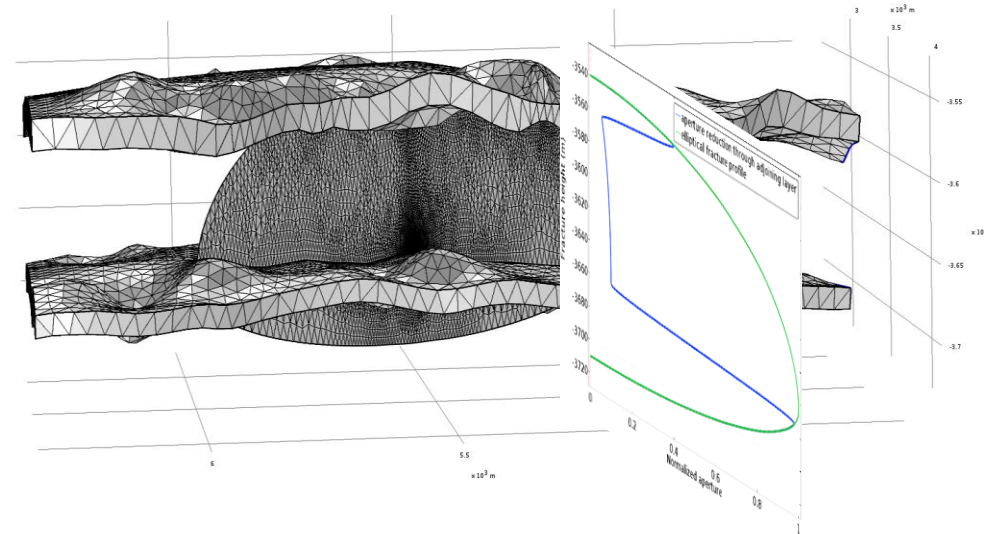
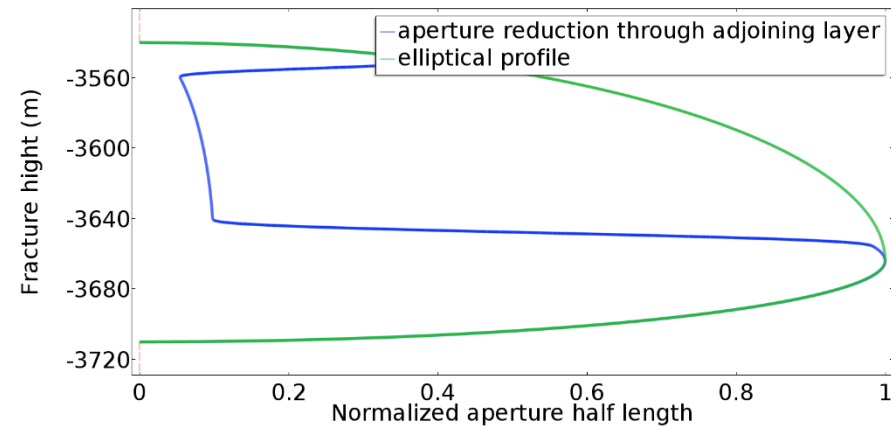
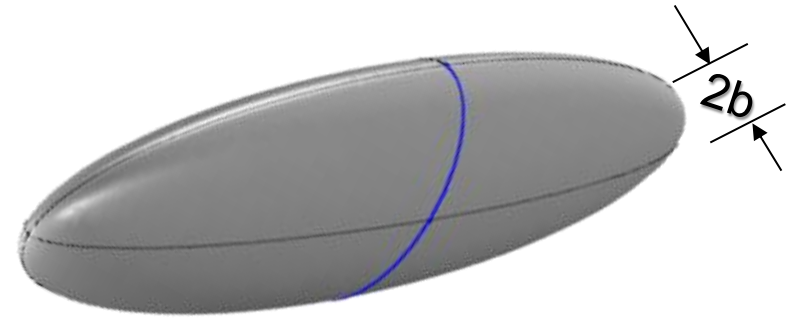


- heat transfer in a fracture is similar to porous media

$$b(\rho C_p)_{eff} \frac{\partial T}{\partial t} + \nabla_t \cdot (b\rho C_p \mathbf{u}_t T) + \nabla_t \cdot \mathbf{q}_{fr} = bQ_h$$

Fracture Dimensions

- Fracture half length : 500 m
- Fracture height : 175 m
- Fracture half aperture (b): 0.0001 m



Base Scenario

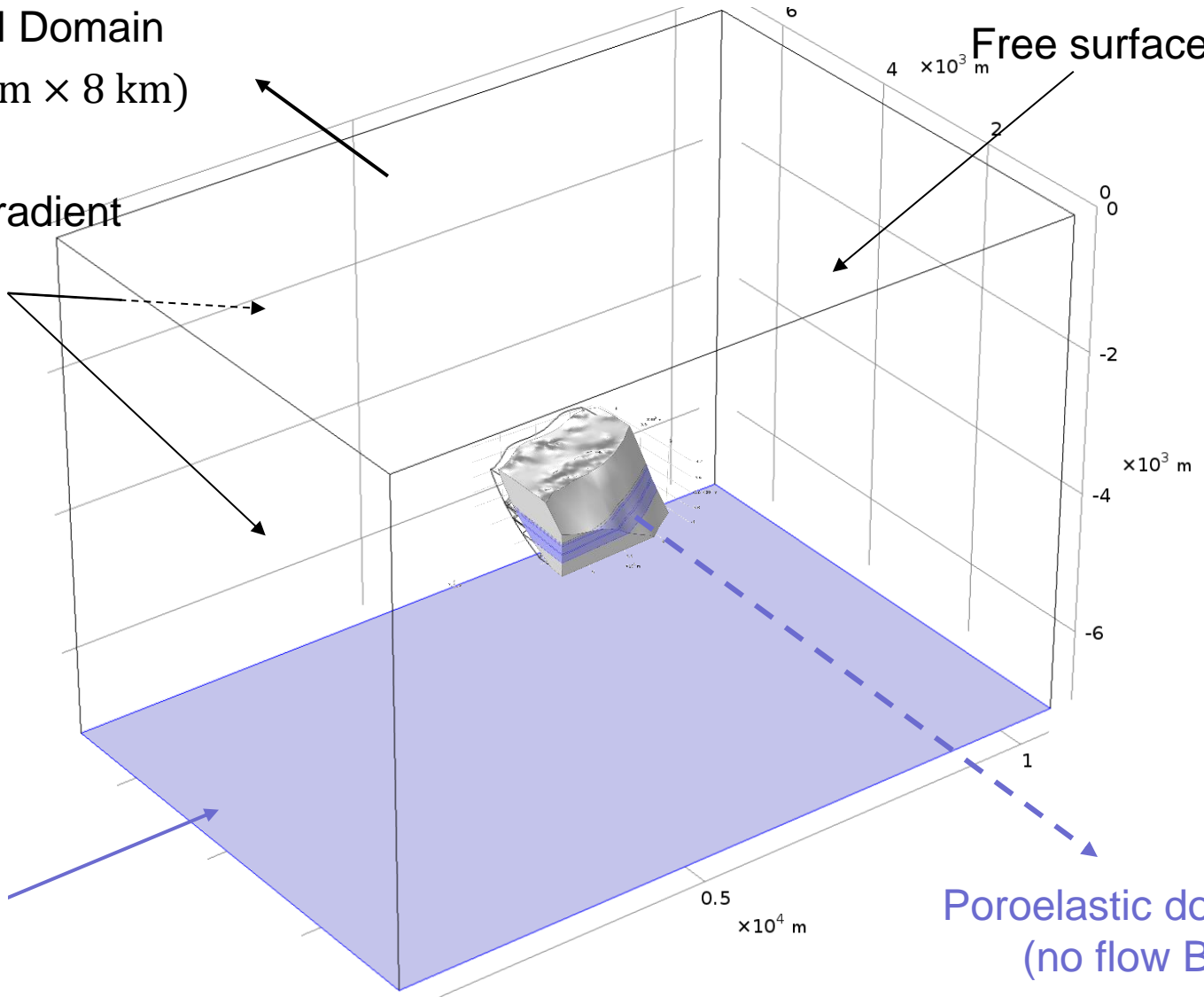
Mechanical Domain
(10 km × 7 km × 8 km)

Lithostatic
pressure gradient

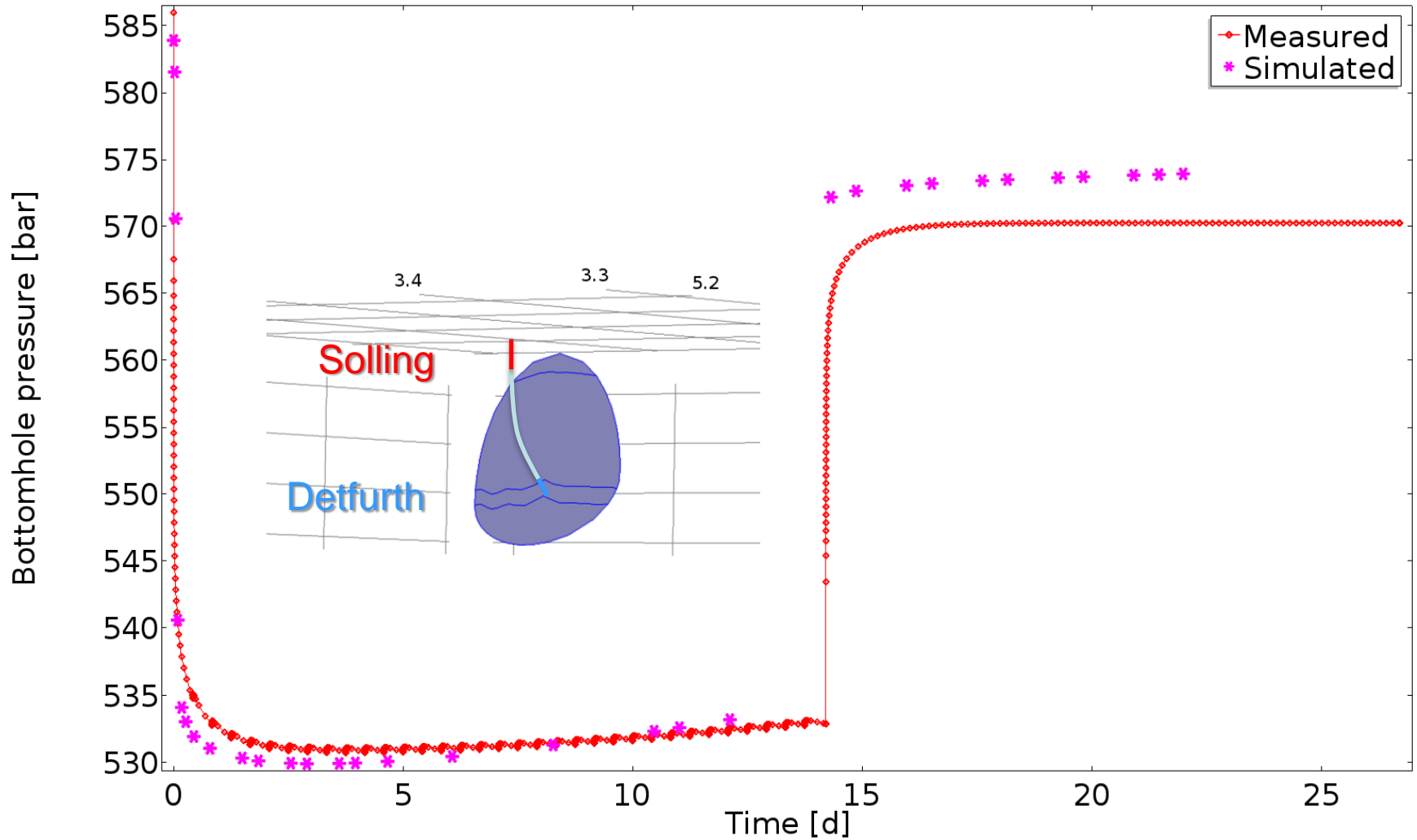
Free surface

Roller

Poroelastic domain
(no flow BC)



Pressure Transient Test - Solling Hydraulic Calibration of the Model



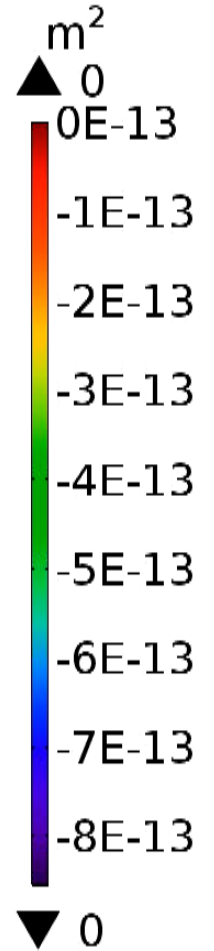
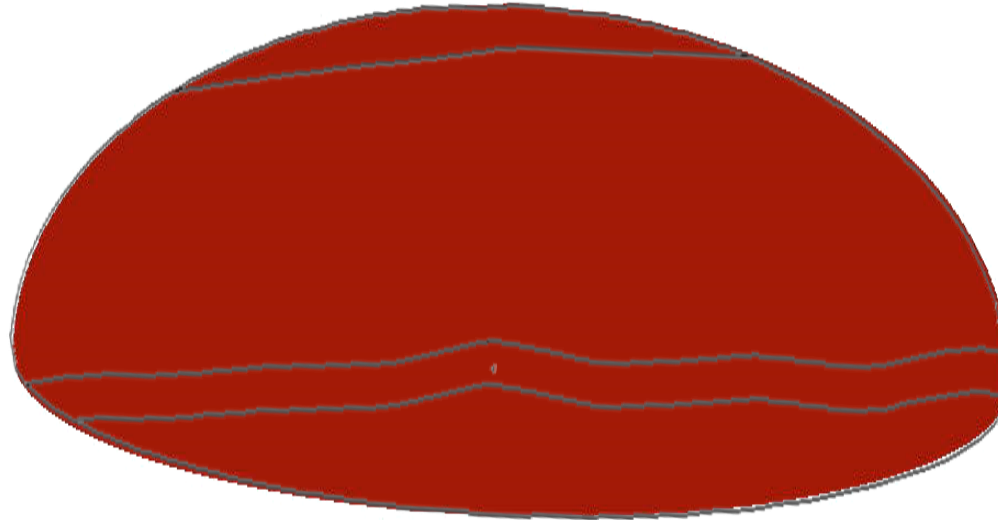
Fracture Permeability Changes

Time=0 d $k_{fr-at}(0, k_{fr})$ (m²)

Solling

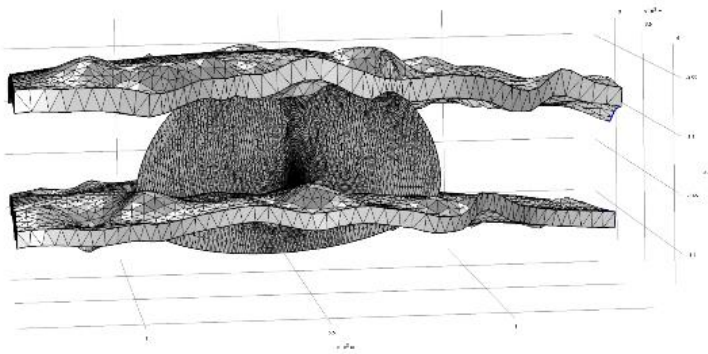
Clay

Detfurth



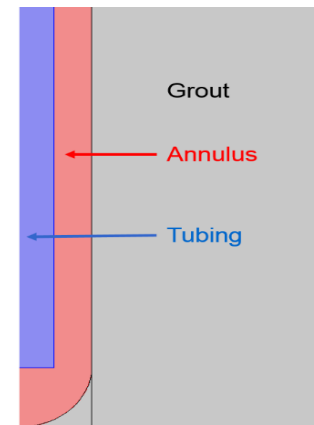
Summary and Outlook

- Comsol Multiphysics is applied to simulate transport phenomena in borehole and reservoir models.
- The coupling aims at transferring data between two models while updating boundary conditions as follow:



Update BCS: →
Production
pressure and
Temperature

←
Update Injection
pressure and
temperature



Thank you, Questions?

I would like acknowledge Wolfram Rühhaak as well as



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